

1 Auxiliary material for

2 **Earthquake clusters in southern California I: Identification and stability**

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10 **Section A. *The earthquake distance: Motivation***

11 The definition (1) of the earthquake distance [Baiesi and Paczuski, 2004] is
12 motivated by the intuitive expectation that the value of η_{ij} should be small if earthquake j
13 might be related to earthquake i , and it should be larger if there is no relationship between
14 earthquakes i and j . To illustrate, consider a situation when $N(m)$ earthquakes with
15 magnitude above m happen independently of each other in d_r -dimensional space and time
16 and obey the Gutenberg-Richter relation $\log_{10}N(m) = a - bm$. Then the expected number of
17 earthquakes with magnitude m within the time interval t and distance r from any given
18 earthquake is proportional to $tr^{d_r}10^{-bm}$, which is an essential component of the definition
19 (1). In other words, the distance (1) is the number (up to a constant) of earthquakes of
20 magnitude m that are expected within the time t and distance r from the earthquake j in a
21 process with no clustering. If the distance η_{ij} is significantly smaller than most pair-wise
22 distance within the catalog, this means that earthquake j has happened abnormally close to
23 i ; this motivates one to consider i as a parent for j . Naturally, this approach only reveals

24 statistical, not causal, relationships between earthquakes. Figure A1 illustrates the
25 connection between the normalized time T (see Eq. (2) of the main text) and the calendar
26 time in years.

27

28 **Section B. *The origin of the bimodal distribution of nearest-neighbor distances***

29 The goal of this section is to shed some light on the origin of the bimodal
30 distribution of the nearest-neighbor distance shown in Fig. 4 of the main text. Comparison
31 of the results for the observed seismicity (Fig. 4) with that for a homogeneous Poisson
32 process (Fig. 3) suggests that the bimodality is related to earthquake clustering. There are
33 several primary types of clustering in the catalogs: time-independent space clustering
34 mainly related to the fault network geometry, space-independent time clustering related to
35 (possible) global changes of seismic activity, and dependent space-time clustering mainly
36 related to the foreshock-aftershock sequences or swarms. We demonstrate below that the
37 cluster mode of the distribution in Fig. 4 *cannot* be explained by temporal or spatial
38 clustering of earthquakes alone. The existence of this mode is ultimately caused by the
39 *clusters* with dependent spatio-temporal structure that are due to the groups of earthquakes
40 that happen within localized spatio-temporal regions; mainly to the foreshock-aftershock
41 sequences or swarms.

42 Towards this goal, we consider three models of seismicity that retain the marginal
43 spatial and/or temporal distributions of the real earthquakes while exhibiting no dependent
44 spatio-temporal clustering. We start with the catalog of observed earthquakes with $m \geq 3$,
45 which contains 12,105 earthquakes. The first randomized catalog is obtained by
46 independent uniform random reshuffling of times and locations of the observed events.

47 Reshuffling means that the event times s_i , $i = 1, \dots, n$, in the new catalog are obtained from
48 the original times t_i , $i = 1, \dots, n$, as $s_i = t_{\sigma(i)}$, where $\sigma(i)$ denotes a uniform random
49 permutation of the sequence $[1, \dots, n]$. An independent reshuffling procedure is then applied
50 to the epicenter locations (ϕ_i, λ_i) . The time-latitude map of seismicity from this catalog is
51 shown in Fig. B1a; the joint distribution (T, R) of the rescaled time and space components
52 of the nearest-neighbor distance is shown in Fig. B2a. By construction, this randomized
53 catalog has the same marginal time and space distributions as the observed seismicity. For
54 instance, in Fig. B1a one can see significant variations of seismic activity along the
55 latitude, which is related to the fault network geometry, as well as the most prominent time
56 variations related to the aftershocks activity in the original catalog. At the same time, we
57 have destroyed all possible clusters with *dependent* spatio-temporal structure. For example,
58 when randomized seismic activity increases in 1992, it affects the entire region, and not
59 only the vicinity of the Landers earthquake as in the original catalog (cf. Fig. 2). Figure
60 B2a shows that this randomization suffices to destroy the bimodal structure of the joint
61 distribution (T, R) : the randomized catalog is characterized by a unimodal distribution of
62 (T, R) located along a diagonal line.

63 The second randomized catalog (Figs. B1b and B2b) is obtained by reshuffling the
64 events locations and using independent uniform random times within the duration of the
65 original catalog. This catalog retains the marginal spatial distribution (and fault-related
66 clustering) of events, while removing all the temporal inhomogeneities. The joint
67 distribution (T, R) is again unimodal; in addition it is more compact and is better separated
68 from the origin, comparing to that of the randomized catalog from Fig. B2a. These
69 differences are related to removing the temporal clustering of events.

70 The third randomized catalog (Figs. B1c and B2c) is obtained by retaining the
71 original times of events and using random locations that are uniformly distributed between
72 30 – 37.5N and 113 – 122W. This catalog retains the temporal clustering of the original
73 catalog while removing all the spatial inhomogeneities. The joint distribution of (T,R) is
74 bimodal in this case, with a weak second mode caused by the temporal clusters. The events
75 that comprise this mode tend to happen close in time to their parents ($T \approx 10^{-6}$) and far
76 away from the parents in space ($R \approx 10^{0.5}$). This spatial separation is two orders of
77 magnitude higher than that observed in the original catalog (Fig. 5b). A noteworthy
78 observation is that the time clustering of the observed seismicity is “stronger” than the
79 spatial clustering, as illustrated by the comparison of the joint distributions (T,R) in Figs.
80 B2b and B2c.

81

82 **Section C. *Proof of the tree structure of the spanning earthquake network***

83 Recall that the NND η is asymmetric: The parent i of event j must happen earlier: t_i
84 $< t_j$. Hence, if we start at any earthquake j in the catalog and repeatedly move from each
85 event to its parent, we never can reach j again. This implies that each possible nearest-
86 neighbor cluster is a tree (a graph without cycles). Next, we show that we only have a
87 single spanning tree. Each nearest-neighbor cluster (tree) must have a *root* – an earthquake
88 without the parent. But we have only one such earthquake – the first event in the catalog;
89 all other events have well-defined parents. This completes the proof.

90

91 **Section D. *Quality and stability of cluster identification in ETAS model***

92

93 **D.1 Model specification and parameters**

94 The ETAS belongs to the class of Marked Point Processes (MPP). Traditionally, the
 95 main object of MPP analysis is the conditional intensity $m(t, \mathbf{f}, m | H_t)$ of a process
 96 $Z_t = \{t_i, \mathbf{f}_i, m_i\}$ given its history $H_t = (\{t_i, \mathbf{f}_i, m_i\} : t_i < t)$ up to time t . Here t_i represents
 97 earthquake occurrence times, \mathbf{f}_i their coordinates (e.g., epicenter, hypocenter, or centroid)
 98 and m_i the magnitudes. It can be shown [Daley and Vere-Jones, 2002] that conditional
 99 intensity completely specifies the process Z_t . The statistical analysis and inference for Z_t are
 100 done using the conditional likelihood

101
$$\log L_t = \sum_{t_i < t} \log \mu(t_i, \mathbf{f}_i, m_i | H_t) - \int_0^t \int_M \int_F \mu(t, \mathbf{f}, m | H_t) dt dm d\mathbf{f}, \quad (D1)$$

102 where M and F denote the magnitude range and spatial domain of events, respectively. The
 103 ETAS assumes a particular self-exciting mechanism of earthquake generation. Namely,
 104 some background events (immigrants) occur according to a homogeneous stationary
 105 Poisson process. Each earthquake in a catalog generates offspring (first generation events),
 106 these offspring generate their own offspring (second generation events), and so on. The
 107 resulting seismic flow is a compound of immigrants and triggered events from all
 108 generations. The main body of the work on ETAS operates under the assumption that the
 109 magnitudes of events are independent and drawn from the Gutenberg-Richter (exponential)
 110 distribution with a constant b -value. This reduces conditional intensity to the following
 111 special form, which allows various particular parameterizations [Ogata, 1998, 1999]:

112
$$\mu(t, \mathbf{f} | H_t) = \mu_0(t, \mathbf{f}) + \sum_{i: t_i < t} g(t - t_i, \mathbf{f} - \mathbf{f}_i, m_i).$$

113 We use in this study a homogeneous background intensity $\mu_0 = \mu$ and the following
 114 parameterization for the response function g suggested by Ogata [1998, Eq. (2.3)]:

115
$$g(t,x,y,m) = \frac{K \exp(\alpha(m - m_0))}{(t+c)^p (x^2 + y^2 + d)^q}. \quad (D2)$$

116 Here m_0 is the lowest considered magnitude, and (x,y) are Cartesian coordinates of the
 117 epicenters. The model is specified by 8 scalar parameters $\theta = \{\mu, b, K, c, p, \alpha, d, q\}$.

118 It has been shown [Sornette and Werner, 2005; Veen and Schoenberg, 2008; Wang
 119 *et al.*, 2010] that estimation of the ETAS model is affected by the catalog’s lowest
 120 magnitude cutoff, which may lead to a serious bias in the numerical values of the estimated
 121 parameters. It is also known that the ETAS parameters depend on the tectonic environment
 122 [Chu *et al.*, 2011] and local physical properties of the lithosphere [Enescu *et al.*, 2009].
 123 These are some of the reasons why there are no commonly accepted “standard” values of
 124 the ETAS parameters for a given region. In this study, we generate synthetic ETAS
 125 catalogs using a range of parameters consistent with those reported in the literature [e.g.,
 126 Wang *et al.*, 2010; Chu *et al.*, 2011; Marzocchi and Zhuang, 2011].

127

128 **D.2 Clustering in ETAS model**

129 An ETAS catalog can be naturally divided into individual clusters according to the
 130 model’s explicit parent-offspring relationships. Namely, a cluster is defined as a group of
 131 events that have the common *ancestor* (grand-parent of arbitrary order), which itself is a
 132 background event (has no parent). This unique cluster’s ancestor is also included in the
 133 cluster; by construction it is always the first event in a cluster. According to this definition,
 134 some clusters consist of a single background event, while the others include several
 135 generation of offspring. Within each cluster, we assign the following event types, same as

136 in analysis of observed catalogs. *Mainshock* is the first largest event in a cluster, *foreshocks*
137 are all events before the mainshock, and *aftershocks* are all events after the mainshock.

138 We next explore how the cluster technique of Sect. 3 can recover (i) the partition of
139 an ETAS catalog into individual clusters, (ii) the event type (main/fore/aftershock)
140 assignment and (iii) the parent-offspring assignment. The analysis is done using the
141 observed catalog of events that reports only their occurrence time, magnitude and location.
142 It should be noted that while we do study the parent-offspring assignment, it plays
143 secondary role in the context of our study, comparing to the partition into individual
144 clusters and event type. In the subsequent analysis, the event types, as well as parent and
145 cluster assignments that correspond to the actual ETAS model structure will be called *true*;
146 while those estimated using our cluster technique will be called *estimated*.

147

148 **D.3 Cluster identification: quality**

149 The analysis in this study was done using multiple ETAS catalogs with a range of
150 realistic parameter values. We found that the results in different catalogs are qualitatively
151 very similar to each other, with quantitative differences being directly related to the model
152 parameters (e.g., different *b*-value, *p*-value, *etc.*) In this and the next section we illustrate
153 the results using a particular ETAS catalog that corresponds to parameters $\mu = 0.003$ (km²
154 year)⁻¹, $b = \alpha = 1$, $K = 0.007$ (km² year)⁻¹, $c = 0.00001$ year, $p = 1.1$, $q = 1.7$, $d = 30$ km²;
155 the simulations are done within a region of 500×500 km during 10 years. The synthetic
156 catalog is illustrated in Figs. D1a, D2a that show, respectively, the magnitude and *X*
157 coordinate of events as a function of time. The catalog consists of 29,761 events, of which
158 7,545 (25%) are background events. Figure D3 shows the joint 2-dimensional distribution

159 of the temporal (T) and spatial (R) components of the nearest-neighbor distance η (panel a)
160 as well as the distribution of the scalar values of η (panel b).The figure clearly
161 demonstrates prominent bimodality of the nearest-neighbor distance, similar to the one
162 reported for the observed seismicity (cf. Fig. 4). A bimodal distribution of the nearest-
163 neighbor distance η in ETAS model has been also reported by *Zaliapin et al. (2008)* and
164 *Gu et al. (2012)*.

165 The time-magnitude and time-coordinate sequence of mainshocks identified by the
166 analyzed cluster technique are illustrated in Figs. D1b and D2b, respectively. Visually, our
167 cluster procedure makes a decent job in identifying and removing the clusters from the
168 original ETAS catalog. Tables D1, D2 and Fig. D4 assess the cluster detection in a
169 quantitative way. Table D1 cross-classifies the events in the catalog according to their true
170 vs. estimated type: 88% of events have been correctly classified into fore/main/aftershocks;
171 the majority of the misclassified events (7%) are aftershocks recognized as mainshocks.
172 The latter misclassification is due to the long-range triggering, when offspring occur at
173 large time and/or distance from their parents. This long-range triggering is caused by the
174 power-law tails of the temporal and spatial offspring kernels use in ETAS model. In the
175 presence of a non-zero background the long-range offspring are mixed with the background
176 events and cannot be correctly identified by a purely statistical procedure; the number of
177 misclassifications increases with the background intensity. Table D2 illustrates similar
178 cross-classification for 279 events with magnitude above 5. Clearly, the quality of detection
179 increases with magnitude of analyzed events. Figure D4 shows the proportion of various
180 misclassifications among events with magnitude above m : Black dots show proportion of
181 events with misspecified parent, open circles – proportion of events assigned to a wrong

182 cluster, squares – proportion of misclassified types (the same as Tables D1, D2), diamonds
183 – proportion of misclassified mainshocks. Notably, the proportion of events with
184 misspecified parents is about 40% for events of magnitude below 6, which is much higher
185 than the proportion of other misclassification types. In particular, the cluster is correctly
186 recognized for over 88% of events; the proportion of respective errors decreases to zero as
187 magnitude m increases to 5.8. This shows that although it can be difficult to detect the true
188 ETAS parents, one can still closely reconstruct the cluster structure of a catalog. This is an
189 important observation, since the clusters present the primary object of the analysis in this
190 study.

191

192 **D.4 Cluster identification: stability**

193 This section assesses and illustrates the stability of cluster identification with
194 respect to the parameters of the algorithm, minimal reported magnitude, catalog
195 incompleteness, and errors in event location.

196 First, we consider the three numerical parameters that are used in the cluster
197 detection procedure: fractal dimension of epicenters d_f , b -value, and cluster detection
198 threshold η_0 . The value of the threshold η_0 is estimated in each experiment from the
199 Gaussian mixture model [Hicks, 2011], except the experiments when we explicitly vary η_0 .

200 We intentionally choose wide ranges for the parameter values:

$$201 \quad 1 \leq d_f \leq 3, 0 \leq b \leq 2, \text{ and } -6 \leq \eta_0 \leq -2.$$

202 The chosen ranges are much wider than the respective statistical margins of error that
203 correspond to estimating these parameters in ETAS model or in observations. This is done
204 in order to test the general limits of applicability of the proposed cluster technique. Recall

205 that the main version of the analysis uses the true ETAS values $d_f = 2$ and $b = 1$ and the
206 corresponding threshold $\eta_0 = -4.476$ from the Gaussian mixture model; we refer to these
207 parameters as *standard*.

208 Figure D5 summarizes the results of 1D stability analysis where we vary a single
209 parameter and keep the rest at their standard values. A rather surprising observation is that
210 the total proportion of misspecified event types, shown in panels (a-c), *never* exceeds 33%,
211 even for obviously outrageous parameter values. For the parameters close to their standard
212 values (shown by stars), the proportion of misspecified events is within 10% – 15%, which
213 is very close to the error of 12% observed in the main version of the analysis. Panel (d)
214 shows individually the proportion of misspecified mainshocks (squares) and aftershocks
215 (triangles) as a function of the threshold η_0 . This panel emphasizes the broadness of the
216 parameter range considered – the proportion of misspecified mainshocks changes from 0 to
217 100% within the considered range. The panel also illustrates that most of the aftershocks
218 are very well separated from the mainshocks: even when the threshold is so low that *all*
219 mainshocks are properly specified, the proportion of misspecified aftershocks is only 40%.
220 The same conclusion can be derived, of course, from visual analysis of the bimodal
221 distribution in Fig. D3.

222 Figure D6 illustrates a 2D stability analysis; it shows the proportions of
223 misspecified mainshocks (panel a) and aftershocks (panel b) as a function of the pair (b, d_f)
224 on a 20x20 grid; the threshold η_0 is estimated in each experiment from a Gaussian mixture
225 model. Similar to the 1D stability experiments, the proportion of errors is a smooth
226 function of the algorithm parameters, so that the error remains close to the one observed for
227 the main version of algorithm. The proportion of misspecified mainshocks in *all*

228 experiments is within 5%-10%. A significant increase of misspecified aftershocks, to 30%,
229 is only observed for clearly “wrong” values of parameters, e.g. $b \approx 0$, $d_f \approx 1$.

230 We now analyze stability of cluster detection with respect to the minimal reported
231 magnitude. Specifically, we perform the cluster analysis for a truncated catalog, only using
232 magnitudes $m \geq m_0$ (starting with computing nearest-neighbor distances, etc.), and then
233 compare the event types estimated in the truncated catalog with the true event types. The
234 results are shown in Fig. D7. The proportion of misspecified events decreases with
235 completeness magnitude m_0 from the original 11.57% to 0 at $m_0 = 5.7$; in other words, the
236 cluster detection quality increases with magnitude of event. The same conclusion can be
237 drawn from the analysis of Fig. D4 above. We notice that the analysis of Fig. D4 differs
238 from the one performed here in that in Fig. D4 we always use the event types estimated in a
239 complete catalog, and only report proportions of errors for different magnitude thresholds.
240 Here, in contrast, we perform the complete cluster and event type estimation in each
241 truncated catalog.

242 Next, we analyze stability of cluster detection with respect to the catalog
243 incompleteness. For that, we perform thinning of the original ETAS catalog so that each
244 event with magnitude $3 \leq m \leq 5$ has probability $P(m) = (5-m)/2$ to be removed. More
245 specifically, all events with magnitude $m \leq 3$ are definitely removed; all events with
246 magnitude $m \geq 5$ are definitely retained; all other events has removal probability $P(m)$ that
247 decreases linearly with magnitude. Figure D8a compares the magnitude distribution in the
248 original and a thinned catalog. The thinning in this experiment is quite severe: it retains
249 only about 20% of events in the catalog. We generate 100 thinned catalogs according to
250 this procedure and compute the proportion of misspecified events in each of them. An

251 event is called misspecified if (i) it has been retained in the catalog after thinning, and (ii)
252 its type in the analysis of the thinned catalog is different from the type of this event in the
253 analysis of the actual catalog. The proportion of misspecified events is 0.1249 ± 0.009
254 (95%CI); its distribution is shown in Fig. D8b. Comparing this with the original
255 misspecification proportion of 0.1157 (see Sect. D3, Table D1), we conclude that the
256 catalog incompleteness has a very weak effect on the cluster detection quality.

257 Finally, we analyze the effects of location errors. For that, we randomly shift the
258 epicenters of events in the ETAS catalog by adding independent 2D Gaussian errors with
259 independent components of zero mean and standard deviation σ . We then perform cluster
260 analysis on a randomized catalog and compare the estimated results with the true ones,
261 focusing on the proportion of the events with misclassified types. We considered 100
262 randomized catalogs for each value of σ . Recall that the cluster identification in the true
263 catalog corresponds to the proportion 0.1157 of misclassified events (see Sect. D3, Table
264 D1). The proportion of misclassified events in randomized catalogs for
265 $\sigma = 0.1\text{km}, 0.3\text{km}, \text{ and } 1.0\text{km}$ is, respectively, 0.1167 ± 0.001 , 0.1170 ± 0.002 , and
266 0.1187 ± 0.002 (95%CI). This shows that random location errors produce practically
267 negligible effect on cluster detection and event classification.

268

269 **D.5 Basic cluster statistics**

270 This section focuses on basic statistics of the detected clusters. The ETAS catalog
271 we use here is longer than the one in the previous sections, to be a better match to the
272 observed catalog in southern California. Specifically, we use an ETAS model with the
273 same parameters as above: $\mu = 0.003 \text{ (km}^2 \text{ year)}^{-1}$, $b = \alpha = 1$, $K = 0.007 \text{ (km}^2 \text{ year)}^{-1}$, $c =$

274 0.00001 year, $p = 1.17$, $q = 1.7$, $d = 30 \text{ km}^2$; the simulations are done within a region of
275 $500 \times 500 \text{ km}$ during 15 years. The catalog consists of 146,432 earthquakes. The bimodal
276 distribution of the nearest-neighbor distance and cluster identification quality (not shown)
277 are similar to those reported in the previous sections for a shorter ETAS catalog.

278 Figure D9 illustrates the frequency-magnitude distribution for mainshocks/singles
279 and aftershocks (true and estimated). The true mainshock and aftershock distributions are
280 distinctly different, each being closely approximated by an exponential (GR) law with
281 different b -values. We also observe upward (downward) deviations from the exponential
282 laws at largest magnitudes. The estimated distributions are very close to the true ones (see
283 legend). Panel (a) shows the cumulative distribution function (cdf), panel (b) shows the
284 normalized cdf in order to emphasize the deviations from a pure exponential law. Table D3
285 reports the maximum likelihood estimations of the b -values for different event types
286 together with the respective uncertainties. A noteworthy observation is that the estimated
287 b -value for aftershocks is larger than that for mainshocks and foreshocks; the same
288 difference is seen in other ETAS catalogs as well (not shown). This difference is due to the
289 conditional assignment of event types, which deflates the b -value for mainshocks (largest
290 events in respective clusters), and, accordingly, inflates it for aftershocks. The b -value for
291 foreshocks is smaller than that for aftershocks since larger events have higher chance to
292 become parents for mainshocks, according to the employed earthquake distance of Eq. (1).

293 Figure D10 illustrates cluster productivity: the number of foreshocks and
294 aftershocks per mainshock. Panel (a) shows the cluster size N as a function of cluster
295 mainshock magnitude m ; the data is closely approximated by the exponential line $N \propto 10^{\beta m}$.
296 The exponent index β estimated within the intermediate magnitude ranges $3.0 \leq m \leq 6.0$ is

297 1.09 ± 0.02 , where the error margins correspond to a 95% confidence interval (95% CI).
298 We also show for comparison the number of first-generation offspring per parent (squares),
299 which by ETAS construction has exponent index 1. Panel (b) shows the cumulative
300 distribution of the cluster size N (circles) and the number of first-generation offspring
301 (squares). Both distributions have a power-law tail. The distribution of the offspring is
302 closely approximated by a Pareto law $F(x) = cx^{-a}$, $c > 0$, $a \approx 1$. The cluster size distribution
303 deviate from this scaling due to finite size effects: The largest events in the catalog tend to
304 attract a larger number of offspring, while the smallest events cannot attract enough
305 offspring because of the catalog's magnitude cutoff. The value of the scaling exponent $a \approx$
306 1 is related to the chosen values of the ETAS parameters $b = \alpha = 1$. It is readily seen (e.g.,
307 *Saichev et al.*, 2005) that the combination of exponential frequency-magnitude relationship
308 with $b = 1$ and exponential offspring productivity with $\alpha = 1$ leads to the power law cluster
309 size distribution with index $a = b/\alpha = 1$. It must be noted though that this argument
310 concerns only the first-generation offspring, while we work with offspring of all
311 generations. We notice, however, that in the examined catalog clusters with only first
312 generation offspring comprise 77% of all non-single clusters, and clusters with the average
313 leaf depth smaller than 2 (hence, with a significant fraction of the first generation offspring)
314 comprise 86% of all non-single clusters. Similar proportions hold for the other examined
315 ETAS catalogs. Hence, the first order approximation to the cluster size distribution can be
316 done under the assumption of single generation offspring.

317 The intensity of foreshocks and aftershocks within 50 days of the mainshock is
318 shown in Fig. D11. Black dots refer to aftershocks (panel a) and foreshocks (panel b) of
319 mainshocks with magnitude $m \geq 4$. The slope of aftershock decay estimated for $t \geq 0.5$ day,

320 is -0.93 ± 0.09 (95%CI); the slope of foreshock decay is harder to estimate due to large
321 fluctuations of the respective intensities. The deviation of the aftershock slope from $p = 1.1$
322 used in ETAS simulations is explained by existence of secondary, ternary, *etc.* aftershocks.
323 Panel (a) shows for comparison (light squares) the intensity of the first-order offspring in
324 ETAS model. The slope estimated within $t \geq 0.5$ day is -1.1 ± 0.01 (95%CI).

325 Figure D12 shows the distribution of magnitude differences between mainshock
326 and aftershock/foreshocks in families with mainshock magnitude $m \geq 4$: panel (a) refers to
327 all aftershocks and foreshocks; panel (b) refers to the largest aftershock/foreshock in a
328 family. The first observation (panel a) is that the majority of aftershocks and foreshocks
329 have rather large magnitude difference from the mainshock: $d_m \geq 4$ for 80% of aftershocks
330 and $d_m \geq 3$ for 80% of foreshocks. It is also noteworthy that the difference Δ_m between the
331 mainshock and the largest aftershock (panel b) is almost uniform within the range $0 \leq \Delta_m \leq$
332 2, while the foreshock difference shows larger fluctuations.

333 Finally, we analyze the distribution of the number N_{off} of direct offspring.
334 According to the ETAS definition, the actual number N_{off} of offspring of an event of
335 magnitude m has Poisson distribution with intensity $\lambda \propto 10^m$. The coefficient of
336 proportionality is determined by the space-time kernel of Eq. (D2). The distribution of the
337 estimated number of offspring though significantly deviates from a pure Poisson. This is
338 explained by the existence of the actual offspring of event i that were attached to other
339 events during the estimation, as well as the offspring of other events that were attached to i .
340 These effects create additional variability in the estimated number N_{off} , which can be
341 closely approximated by a *negative binomial* distribution, as illustrated in Fig. 12b of the
342 main text.

343

344 **Section E. *Stability of cluster identification in southern California***

345 This section assesses the stability of cluster identification in the observed catalog.
346 Here, unlike the analysis of ETAS model, we do not know the “true” cluster structure, so
347 the *quality* of cluster identification cannot be directly assessed. At the same time, we can
348 assess its *stability*. For that, we vary parameters of the algorithm and compare results with
349 the ones obtained in the main version of the analysis, which is done here with $d_f = 1.6$, $b =$
350 1 , minimal magnitude $m_0 = 3$, and threshold η_0 estimated from the Gaussian mixture
351 model. The use of adaptive estimation of the threshold is important in these experiments,
352 since its values depend (although weakly) on the other three parameters of the algorithm.
353 Figure E1 shows the proportion of events with estimated type different from that obtained
354 in the main version of analysis, as a function of each of the parameters. Similarly to the
355 ETAS stability analysis, we intentionally use very wide ranges for parameter variation, in
356 order to explore the general limits of algorithm stability:

357
$$1 \leq d_f \leq 2, 0 \leq b \leq 2, 3 \leq m_0 \leq 6, \text{ and } -6 \leq \eta_0 \leq -4.$$

358 The proportion of misspecified types is below 7% for *all* experiments within the following
359 parameter ranges:

360
$$1.1 \leq d_f \leq 2, 0.5 \leq b \leq 1.3, 3 \leq m_0 \leq 6 \text{ and } -5.5 \leq \eta_0 \leq -4.55.$$

361 The errors larger than 7% are only observed for the parameter values that are clearly
362 inconsistent with the available observations, like $b > 1.5$. Notably, the proportion of errors
363 *never* exceeds 18% in our experiments.

364 Next, we analyze the stability of cluster detection with respect to the event location
365 error. Specifically, we generate 100 catalogs by randomly altering the locations of events.

366 The location error is modeled by a 2D Normal random variable with zero mean,
367 independent components, and standard deviation for both component given by the standard
368 error of event location reported by *Hauksson et al.* (2012). The proportion of misspecified
369 event types (compared to the analysis of true event locations) is 0.044 ± 0.005 (95% CI); the
370 maximal observed proportion is 0.051. This shows that the proposed algorithm is stable
371 with respect to the location uncertainties.

372 The *stability* results of this section are consistent with that obtained above in ETAS
373 model. This supports a conjecture that the *quality* of cluster detection, if one assumes that
374 there exists a *true* cluster structure in observed catalogs, is also good, similar to that in
375 ETAS analysis.

376

377 Table D1: Cross-classification of event types (true vs. estimated) in ETAS catalog:

378 All 29,671 events are considered

		True		
		Foreshock	Mainshock	Aftershock
Estimated	Foreshock	2760 (9%)	77 (0.2%)	157 (0.5%)
	Mainshock	331 (1%)	7007 (24%)	2198 (7%)
	Aftershock	242 (0.8%)	461 (2%)	16438 (55%)

379

380

381 Table D2: Cross-classification of event types (true vs. estimated) in ETAS catalog:

382 279 events with magnitude $m \geq 5$ are considered

		True		
		Foreshock	Mainshock	Aftershock
Estimated	Foreshock	31 (11%)	1 (0.4%)	1 (0.4%)
	Mainshock	6 (2%)	90 (32%)	11 (4%)
	Aftershock	-	4 (1%)	135 (48%)

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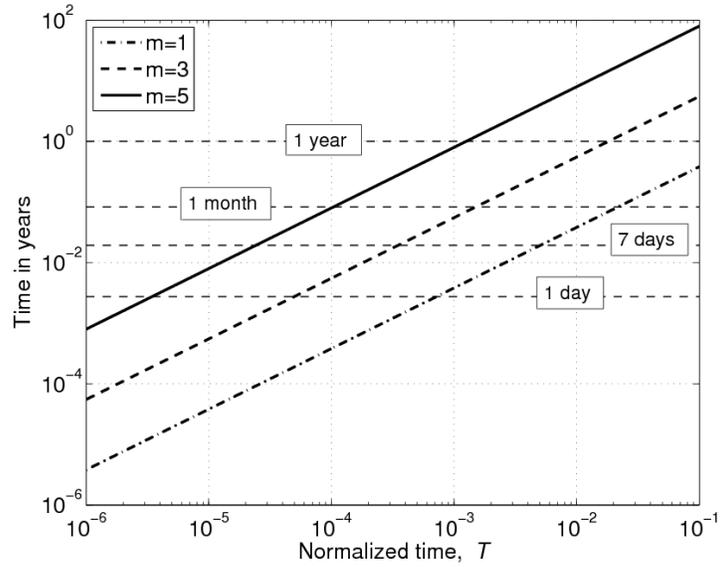
Table D3: Estimated b -values for different event types in ETAS catalog

391

(maximum likelihood estimation and confidence interval)

	True		Estimated	
	b -value	95% CI	b -value	95% CI
Mainshocks	0.932	0.91 – 0.95	0.957	0.94 – 0.97
Aftershocks	1.006	1.00 – 1.01	1.006	1.00 – 1.01
Foreshocks	0.960	0.92– 1.00	0.935	0.89 – 0.98

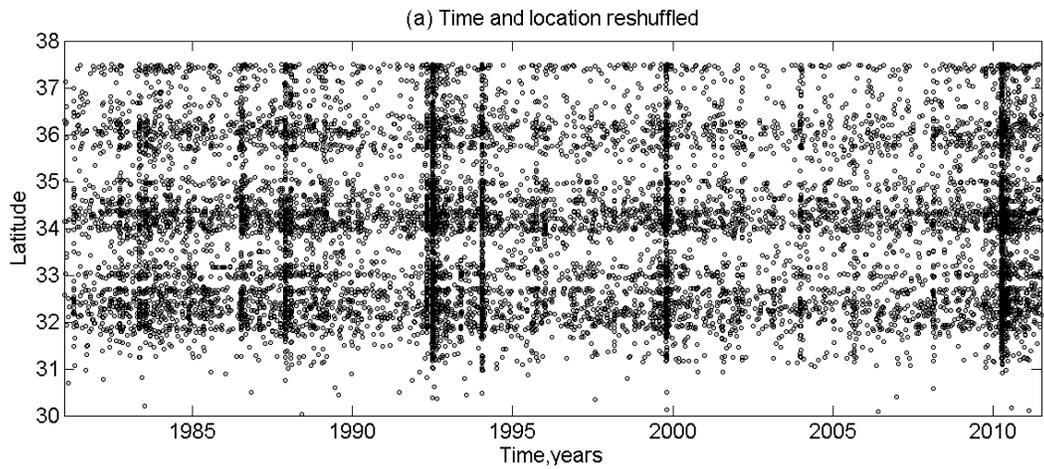
392



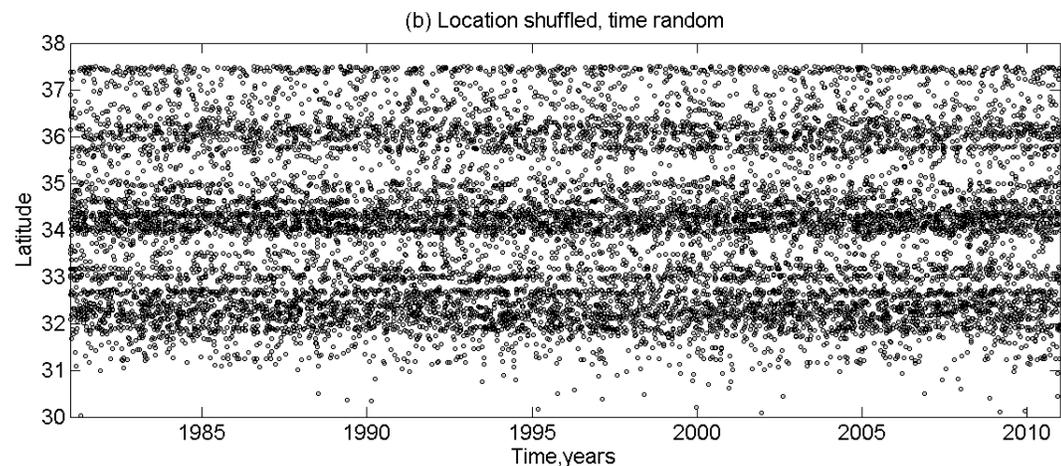
393

394 Figure A1: Correspondence between the normalized time T of Eq. (2) (x -axis) used in the
 395 2-D cluster analysis and time in years (y -axis) for earthquakes of different parent
 396 magnitudes, $m = 1, 3, \text{ and } 5$. Horizontal lines indicate times of 1 day, 7 days, 1 month, and
 397 1 year.

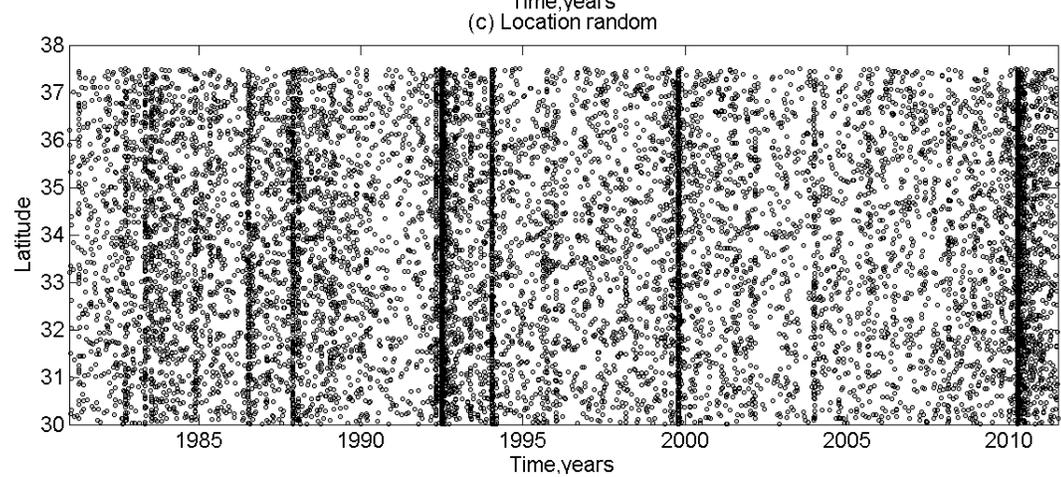
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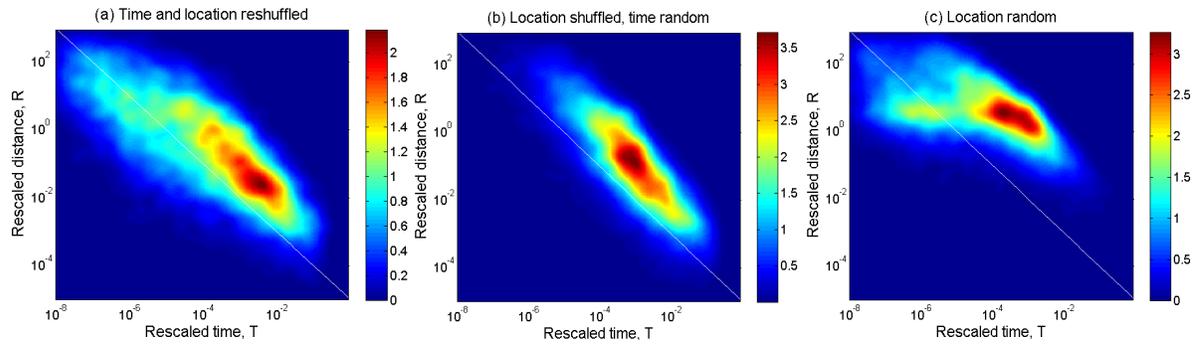


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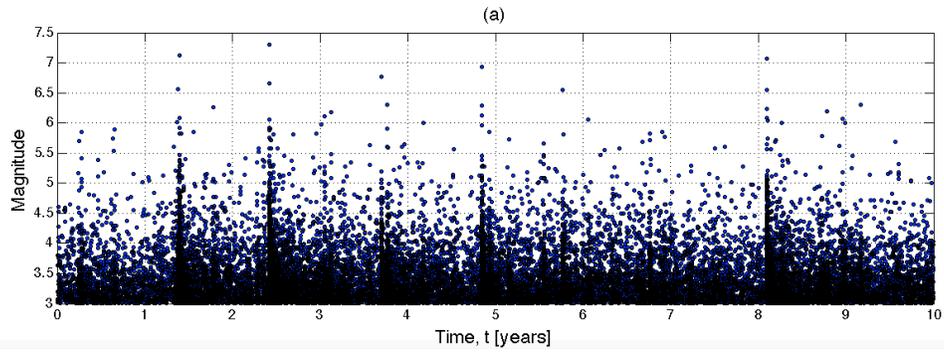
402 Figure B1: Time-latitude map of earthquakes from randomized catalogs. (a) Times and
 403 locations of the observed events are randomly reshuffled. (b) Locations are randomly
 404 reshuffled; times are uniform random variables. (c) Locations are uniform random
 405 variables, original times.



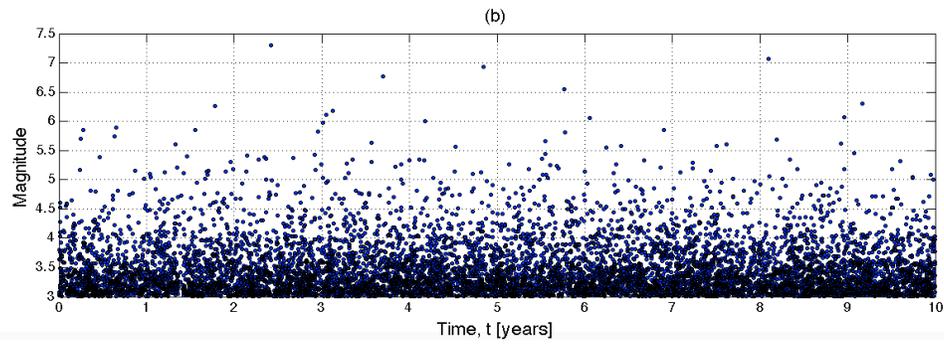
406

407 Figure B2: The joint distribution of rescaled time and space components (T,R) of the
 408 nearest-neighbor distance η in randomized catalogs. (a) Times and locations are randomly
 409 reshuffled. This catalog retains the marginal spatial and temporal distributions of the
 410 observed seismicity, while removing their local interactions. (b) Locations are randomly
 411 reshuffled; times are uniform random variables. This catalog retains the spatial clustering,
 412 while removing all the time inhomogeneities. (c) Locations are uniform random variables,
 413 original times. This catalog retains the temporal clustering, while removing all the space
 414 inhomogeneities.

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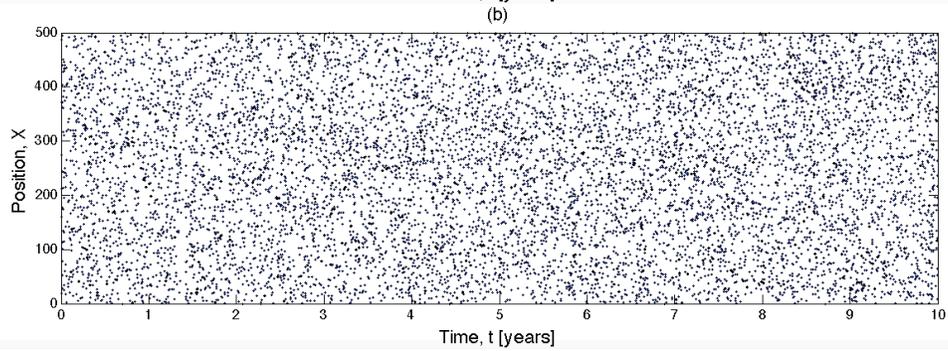
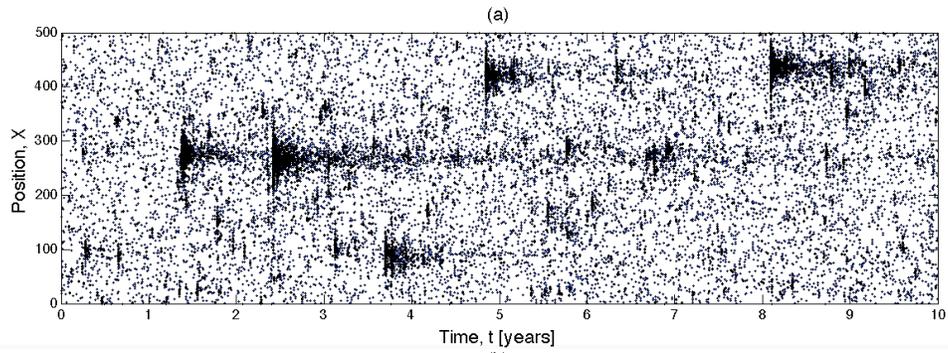
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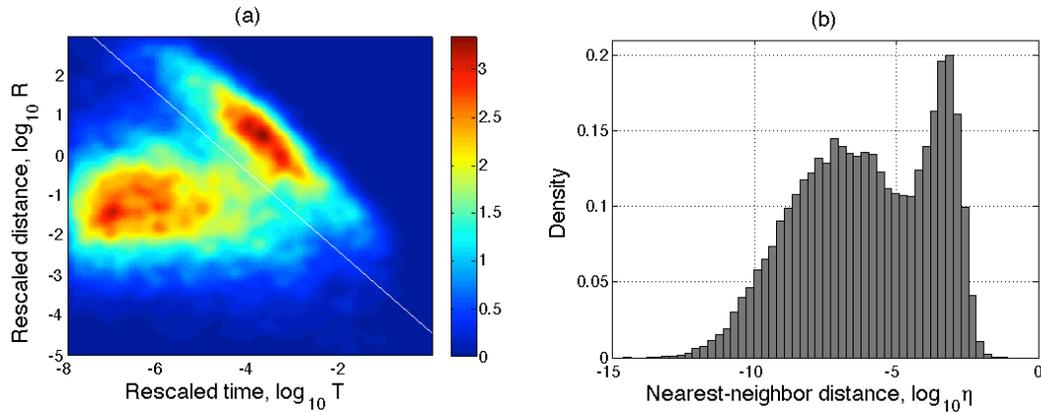
418 Figure D1: ETAS model – an example of declustering. Figure shows the time-magnitude
419 sequence for events with $m \geq 3$. (a) All events, $n = 29,671$; (b) Mainshocks, $n = 9,536$.

420



423 Figure D2: ETAS model – an example of declustering. Figure shows the X coordinate of
424 epicenters vs. time for all events in the catalog. (a) All events, $n = 29,671$; (b) Mainshocks,
425 $n = 9,536$.

426



427

428 Figure D3: ETAS model – nearest-neighbor distance. (a) Joint distribution of the time and

429 space components (T, R) of the nearest-neighbor distance η . (b) Histogram of the log-values

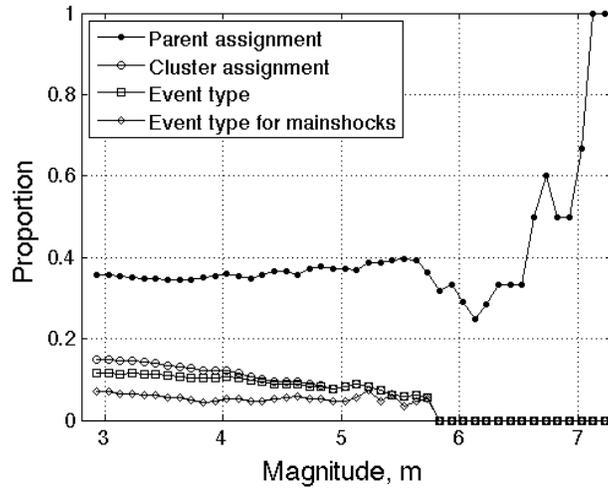
430 of the nearest-neighbor distance η . Bimodal distribution is clearly seen: the background

431 part is located above the white line in panel (a), and corresponds right mode in panel (b);

432 clustered part is located below the white line in panel (b), and corresponds to left mode in

433 panel (b). The white line in panel (a) corresponds to $\eta = -4.47$.

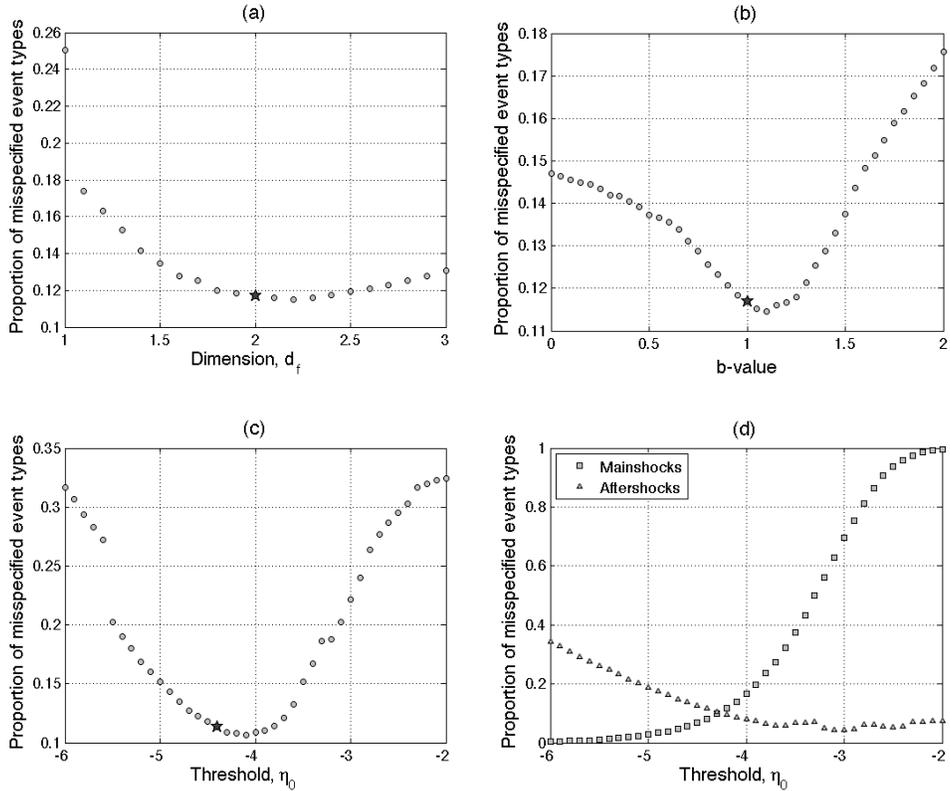
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435

436 Figure D4: ETAS model – cluster identification errors. The figure shows the proportion of
 437 various erroneous identifications for events with magnitude above m . Dots – wrong parent
 438 assignment; circles – wrong cluster assignment; squares – wrong event type
 439 (fore/after/mainshock) assignment, stars – wrong event type assignment for mainshocks
 440 only.

441

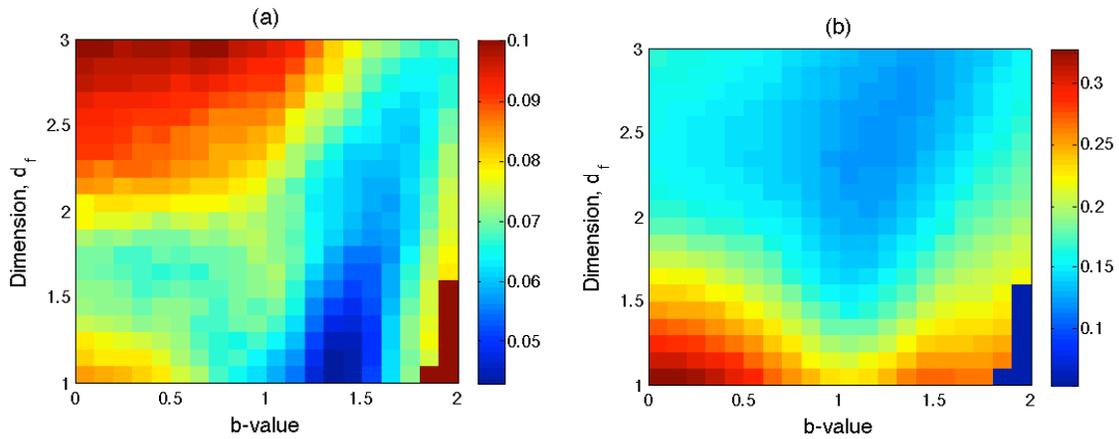


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443

444 Figure D5: ETAS model – stability of cluster identification. Proportion of events with
 445 misspecified event type vs. model numerical parameters. Each panel refers to variation of a
 446 single parameter with the other parameters fixed. Stars in panels (a)-(c) refer to the values
 447 that correspond to the main version of the analysis, with true values of $d_f = 2$, and $b = 1$,
 448 and η_0 estimated according to the Gaussian mixture model. See text for details.
 449 Specifically, we vary (a) the fractal dimension d_f of epicenters, (b) b -value, and (c-d) the
 450 threshold η_0 . Panels (a-c) show the proportion of all events with misspecified type, panel
 451 (d) shows separately the proportion of misspecified mainshocks (squares) and aftershocks
 452 (triangles).

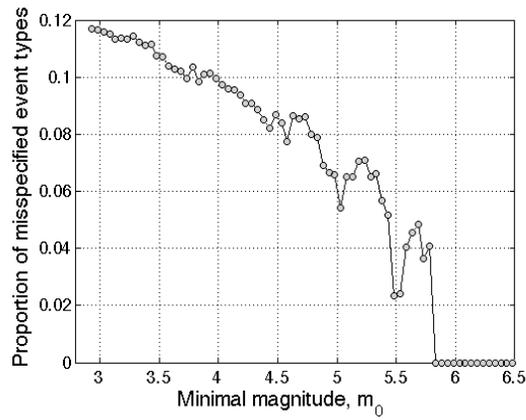
453



454

455 Figure D6: ETAS model – stability of cluster identification. Proportion of misspecified
 456 mainshocks (panel a) and aftershocks (panel b) as a function of the pair (b, d_f) .

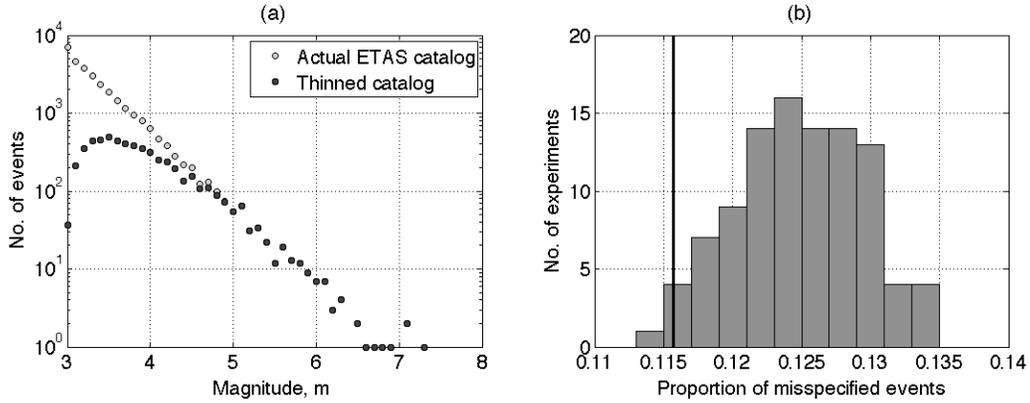
457



458

459 Figure D7: ETAS model – stability of cluster identification. Proportion of events with
 460 misspecified types, as a function of minimal magnitude of analysis.

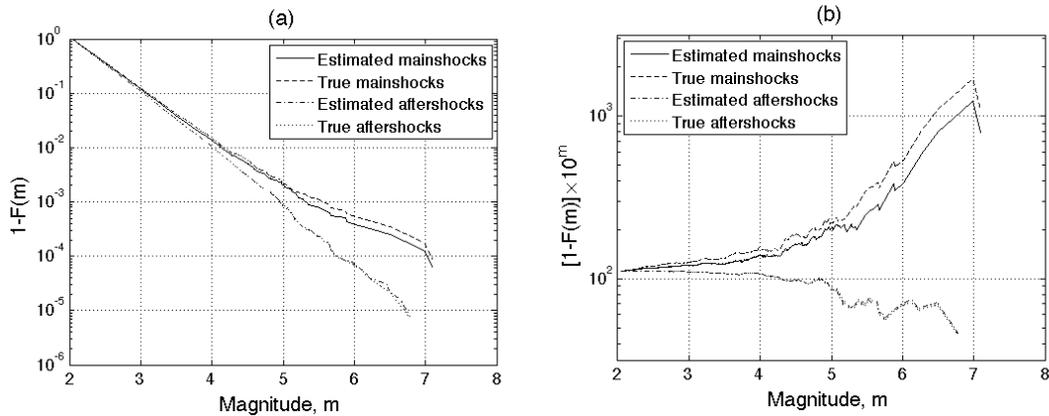
461



462

463 Figure D8: ETAS model – stability of cluster identification in thinning experiment. A
 464 thinned catalog is obtained from the actual catalog by removing each event with probability
 465 $P(m)$ that decrease linearly from 1 to 0 on the interval $3 \leq m \leq 5$. (a) Magnitude distribution
 466 in the actual (black circles) and a thinned (light circles) catalog. (b) Distribution of the
 467 proportion of misspecified events for 100 thinned catalogs. Black vertical line refers to the
 468 proportion of misspecified events in the true, complete catalog.

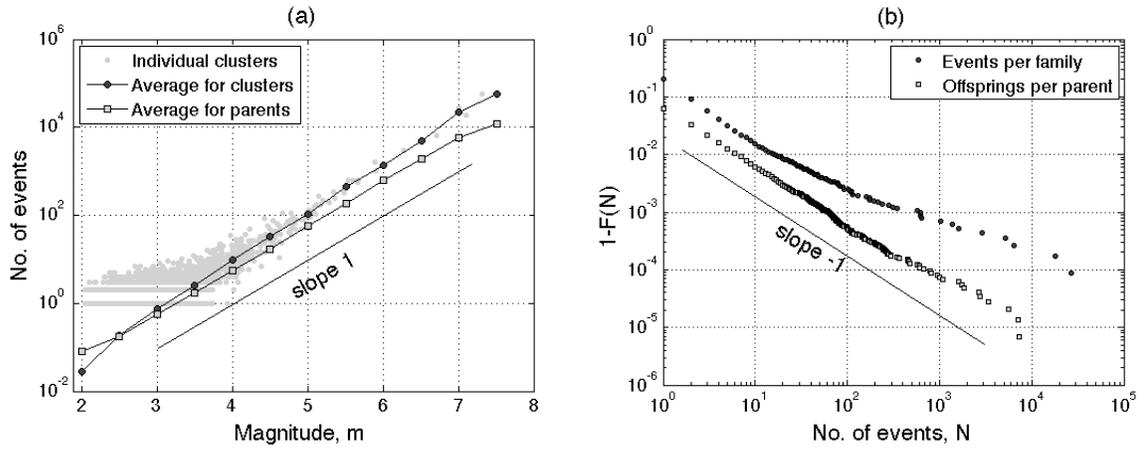
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470

471 Figure D9: ETAS model – magnitude-frequency distribution. Figure refers to different
 472 event types as described in the legend. (a) Proportion $1-F(m)$ of events with magnitude
 473 above m , where $F(m)$ is the empirical cumulative distribution function. (b) Weighted
 474 proportion of events with magnitude above m , $[1-F(m)] \times 10^m$. Panel (b) emphasizes

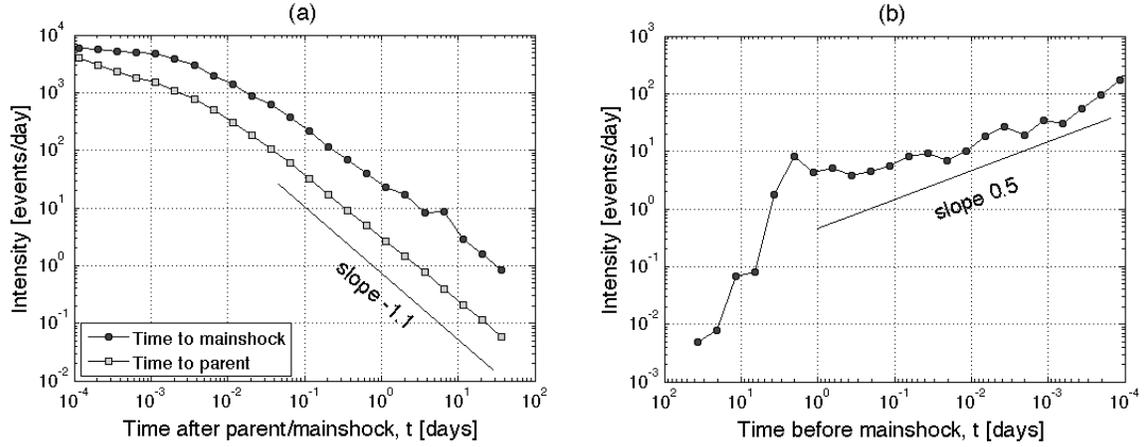
475 deviations from an exponential distribution $E(m) = 1-10^{-m}$ with b -value 1, which
 476 corresponds to a horizontal line.



477

478 Figure D10: ETAS model – cluster productivity. (a) Number of aftershocks and foreshocks,
 479 $N-1$, in a cluster vs. cluster magnitude m . Black circles – average number of events in a
 480 cluster within magnitude window of length 0.5. Grey dots – individual clusters. Squares –
 481 average number of offspring per parent. (b) Distribution of cluster size N (black circles)
 482 and the number of offspring per parent (squares).

483



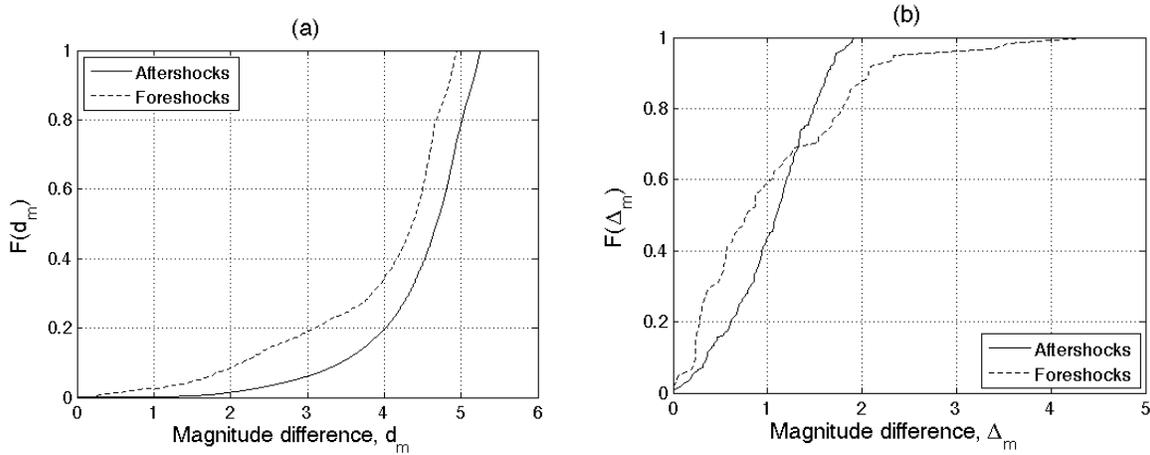
484

485 Figure D11: ETAS model – Aftershock and foreshock intensity. (a) Black dots –

486 aftershocks within 50 days of mainshocks with magnitude $m \geq 4$. Squares – first generation

487 offspring. (b) Foreshocks within 50 days of mainshocks with magnitude $m \geq 4$.

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489

490 Figure D12: ETAS model – magnitude difference analysis. (a) Magnitude difference d_m

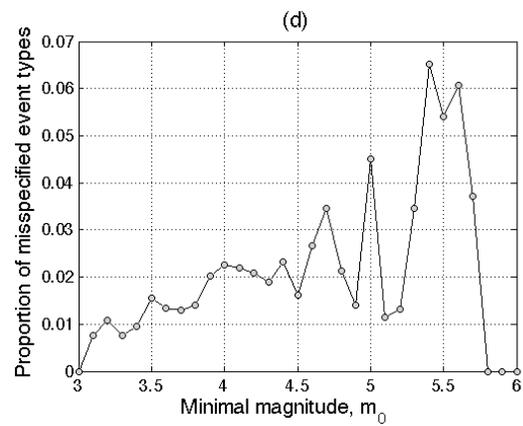
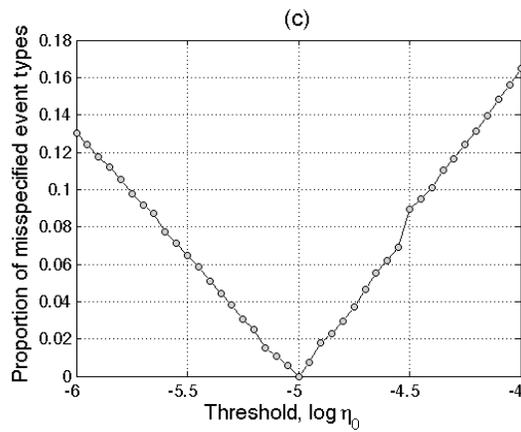
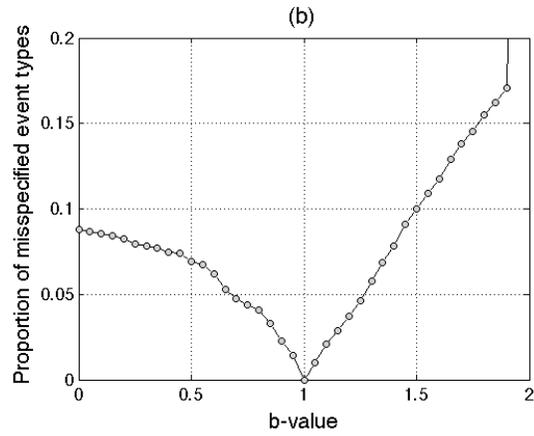
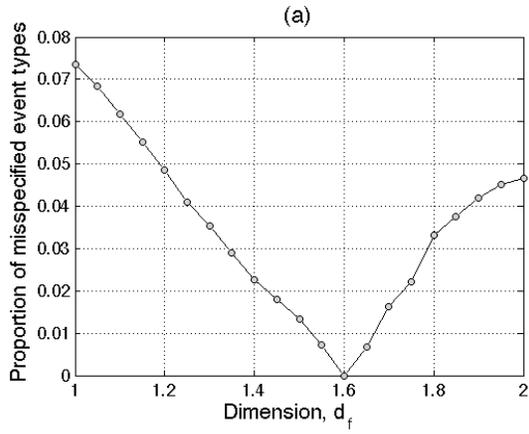
491 between mainshock and each aftershock (solid line) and foreshock (dashed line). (b)

492 Magnitude difference Δ_m between mainshock and the largest aftershock (solid line) and

493 largest foreshock (dashed line). Families with mainshock magnitude $m \geq 4$ are considered

494 in both panels.

495



496

497

498 Figure E1: Stability of cluster identification in southern California. Proportion of events
 499 with event type different from that obtained in the main version of analysis as a function of
 500 algorithm parameter: (a) Fractal dimension of epicenters d_f , (b) b -value, (c) cluster
 501 threshold η_0 , and (d) minimal magnitude of analysis. The main version of analysis uses $d_f =$
 502 1.6 , $b = 1$, $m_0 = 3$, and threshold η_0 estimated from the Gaussian mixture model.